

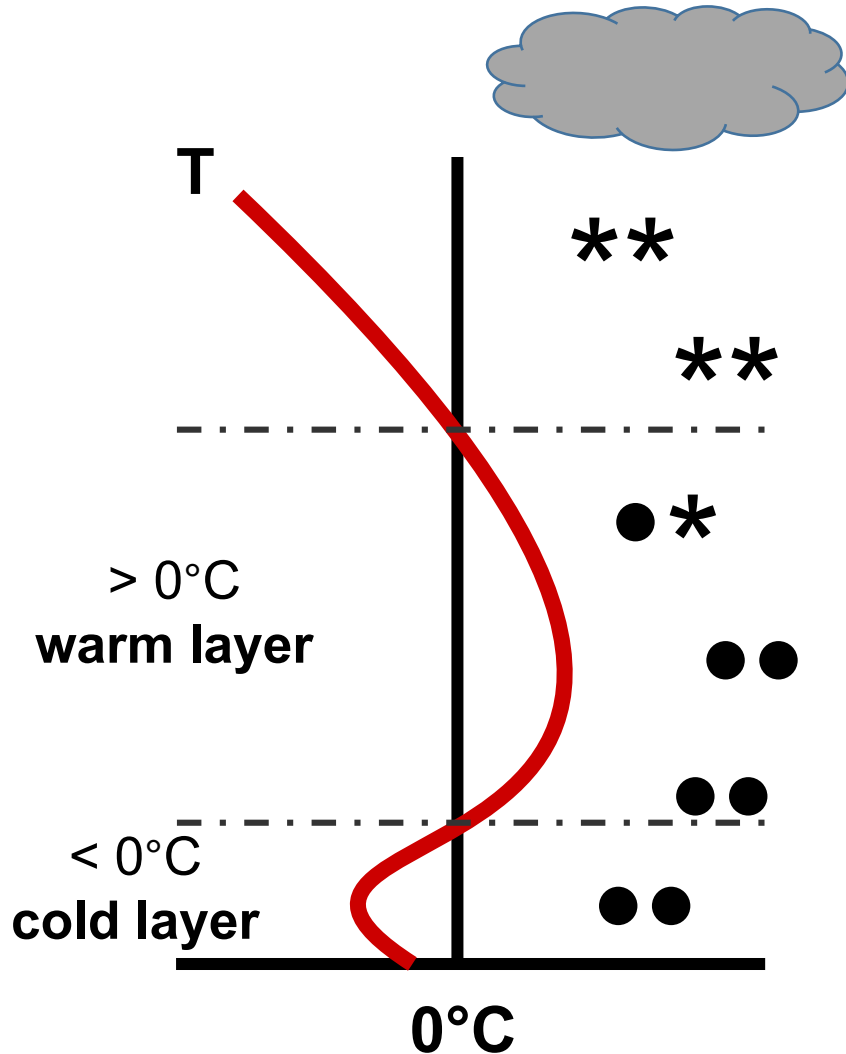
Thermodynamic and Synoptic-Dynamic Modulations of Freezing Rain Event Duration

Christopher McCray, John Gyakum and Eyad Atallah
McGill University - Dept. of Atmospheric and Oceanic Sciences

44th Northeastern Storm Conference - 9 March 2019

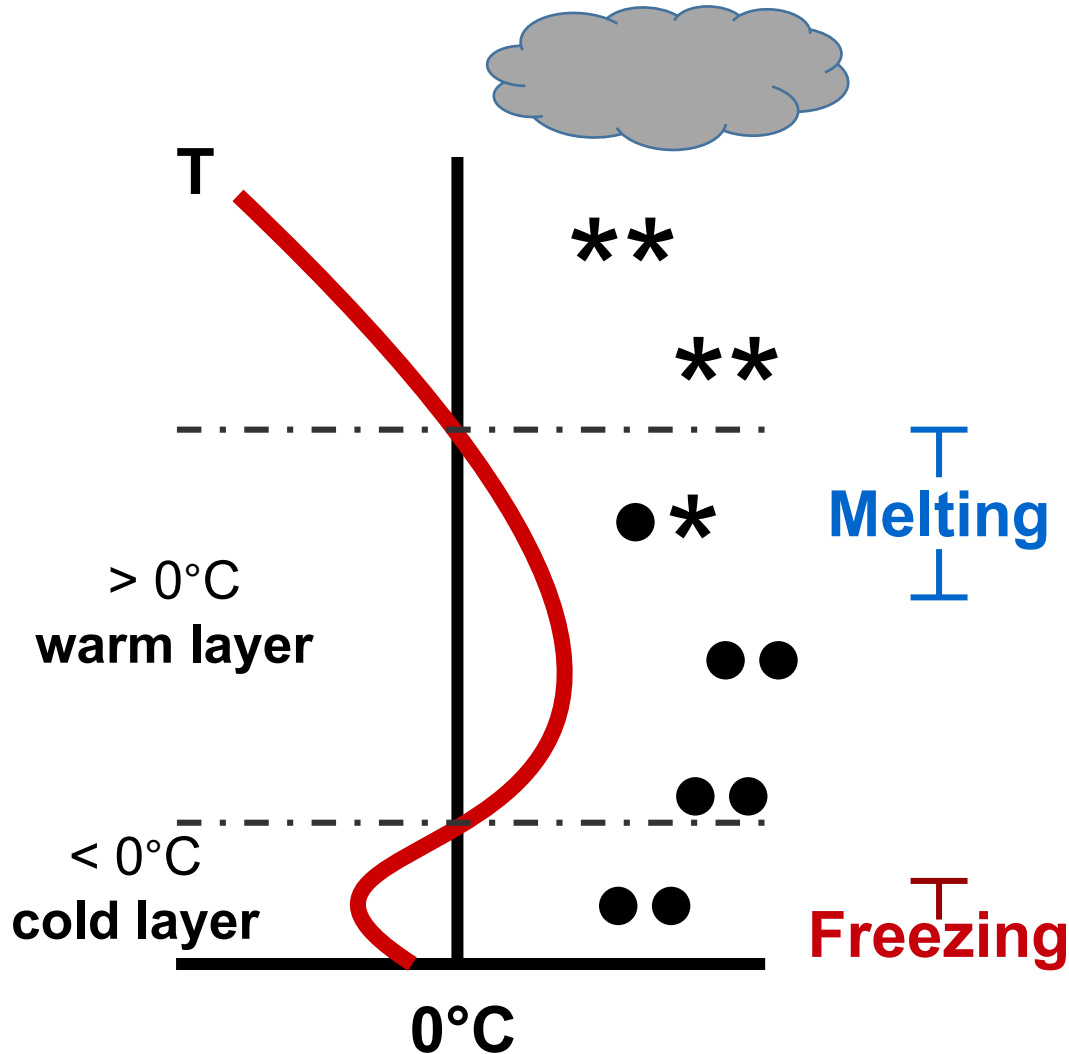


Freezing rain events remain a major forecast challenge



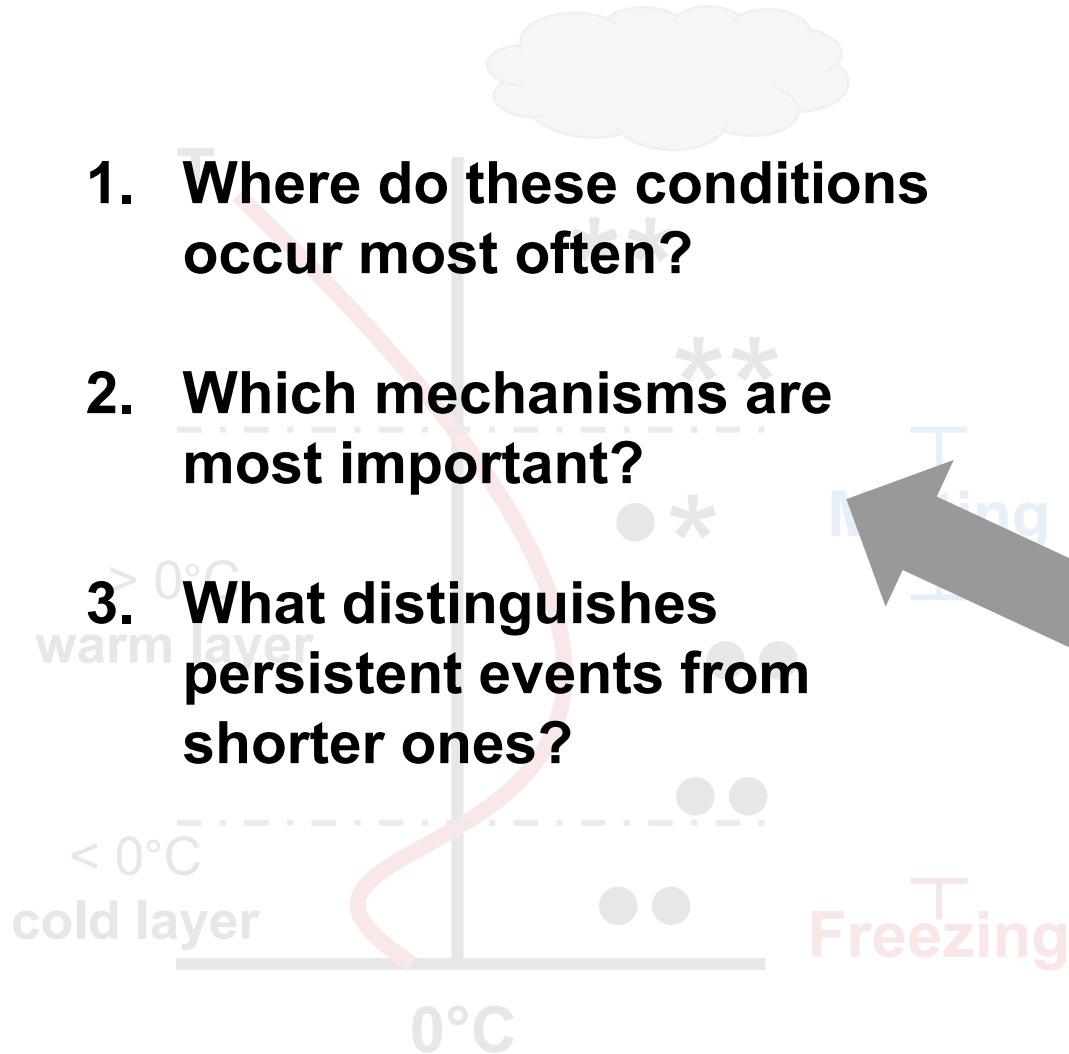
- They require a sufficiently warm **warm layer** aloft to completely melt snowflakes
- They also require **surface temperatures $< 0^{\circ}\text{C}$**
- Only **11%** of FZRA events last longer than 4 h (Cortinas et al. 2004)
 - **Duration** is a key factor in event severity

Diabatic effects mean freezing rain events are *self-limiting* (Stewart et al. 1985, Lackmann et al. 2001)



- Latent heat of fusion is...
 - **Extracted** when snowflakes melt in warm layer (cooling the warm layer)
 - **Released** when rain freezes at the surface (warming the cold layer)
- For event persistence, compensatory mechanisms are necessary:
 - At the surface...
 - Particularly **cold, dry onset conditions**
OR **advection of cold, dry air**
 - In the warm layer
 - Particularly **warm onset conditions**
OR **warm-air advection**

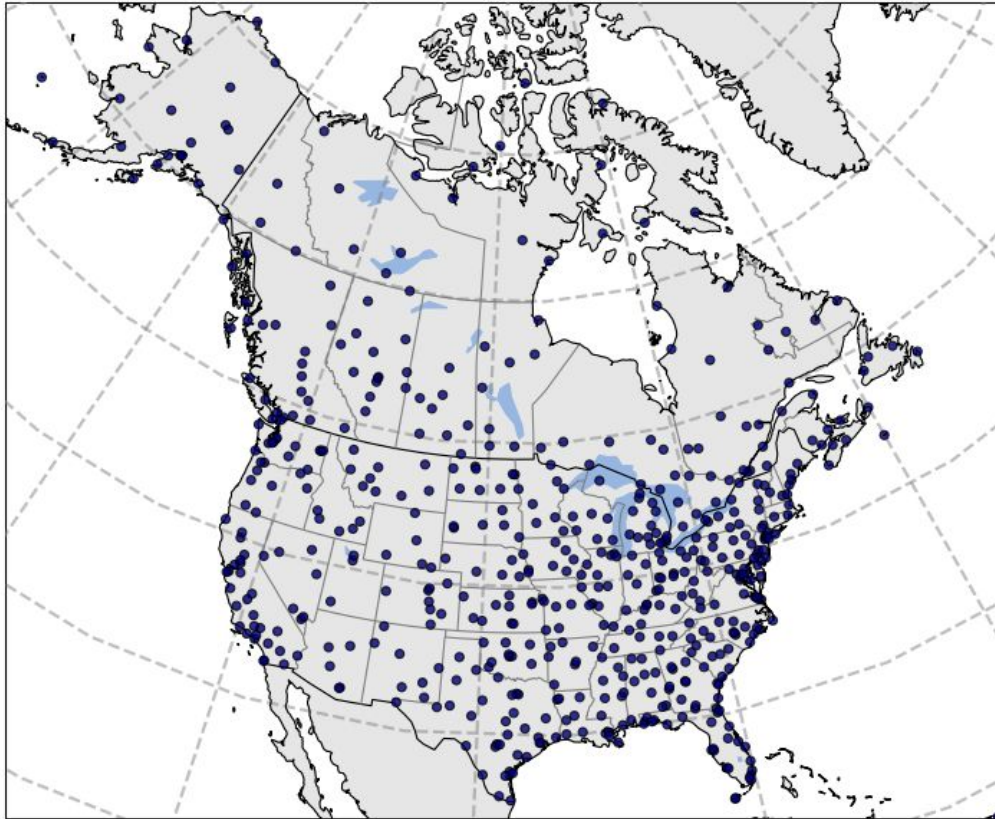
Diabatic effects mean freezing rain events are *self-limiting* (Stewart et al. 1985, Lackmann et al. 2001)



- Latent heat of fusion is...
 - **Extracted** when snowflakes melt in warm layer (cooling the warm layer)
 - **Released** when rain freezes at the surface (warming the cold layer)
- For event persistence, compensatory mechanisms are necessary:
 - At the surface...
 - Particularly **cold, dry onset conditions**
OR **advection of cold, dry air**
 - In the warm layer
 - Particularly **warm onset conditions**
OR **warm-air advection**

Data and methods

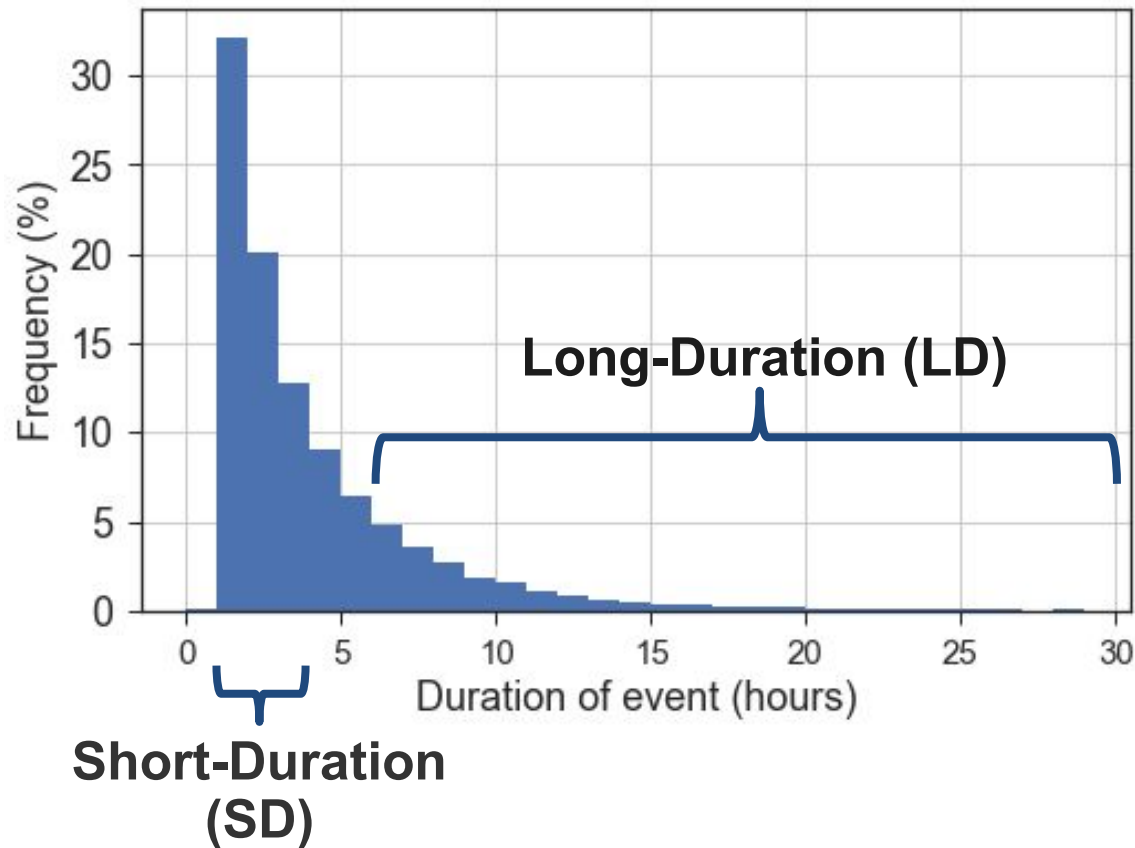
579 surface stations used in dataset



- **Surface Observations:**
 - NOAA Integrated Surface Database
 - 1979-2016, U.S. and Canada
- **Upper-air data:** U. Wyoming archive
- **NCEP CFSR Reanalysis**
 - 0.5°x0.5° grid, 6-hourly, 1979-present

Data and methods

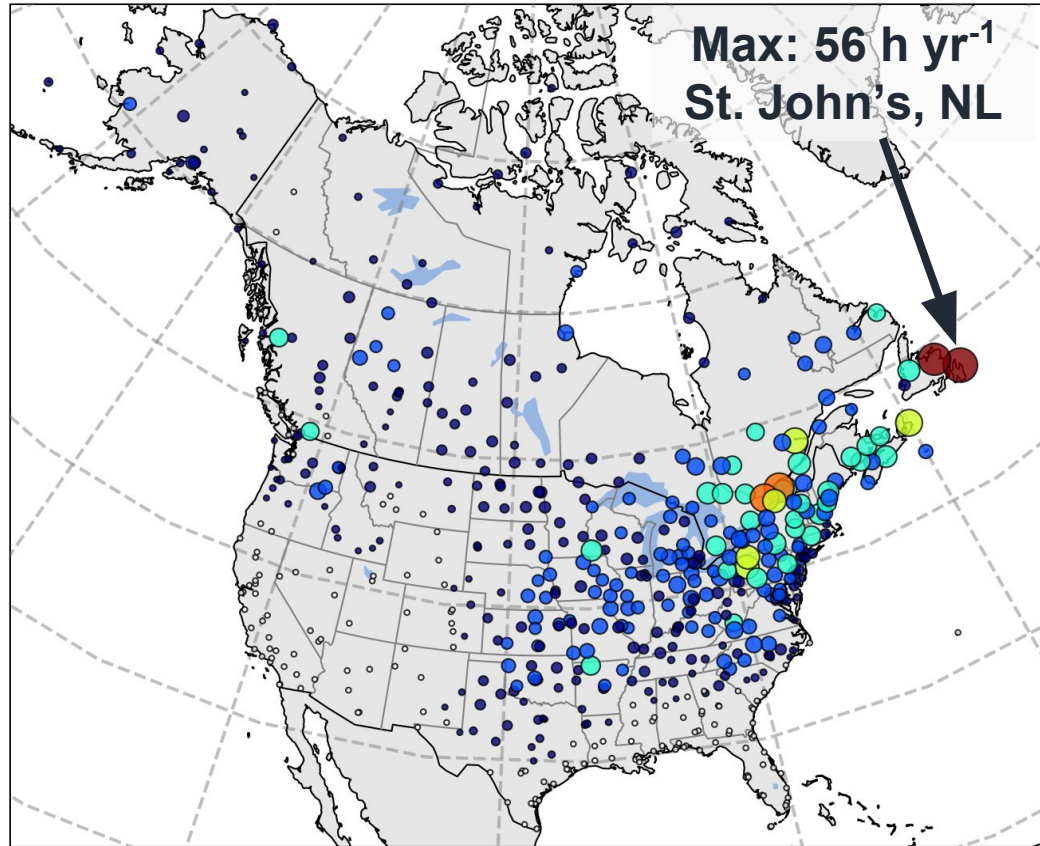
Histogram - Event Duration



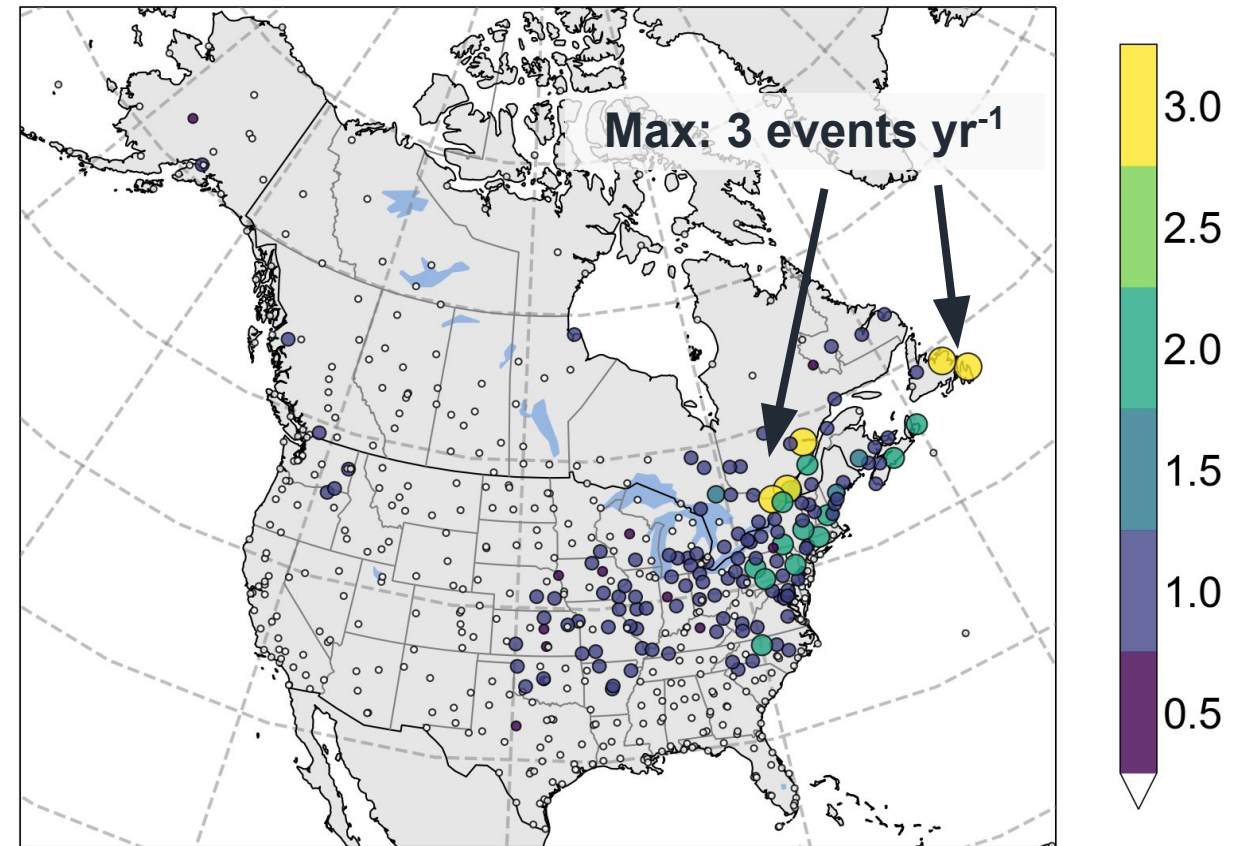
- **Surface Observations:**
 - NOAA Integrated Surface Database
 - 1979-2016, U.S. and Canada
- **Upper-air data:** U. Wyoming archive
- **NCEP CFSR Reanalysis**
 - 0.5°x0.5° grid, 6-hourly, 1979-present
- **Freezing Rain Event Duration:**
 - Count consecutive hours of FZRA, then combine events with <24 h between them
- **Long-Duration (LD) Event:**
 - FZRA event with **6+ h of FZRA**
 - ~20% of all events
 - Longest: 61 h (Montréal, 1998 Ice Storm)
- **Short-Duration (SD) Event:**
 - FZRA event with **≤ 3 h of FZRA**

Freezing rain, LD events occur most often in the northeastern U.S. and southeastern Canada...

Median Annual FZRA Hours (1979-2016)

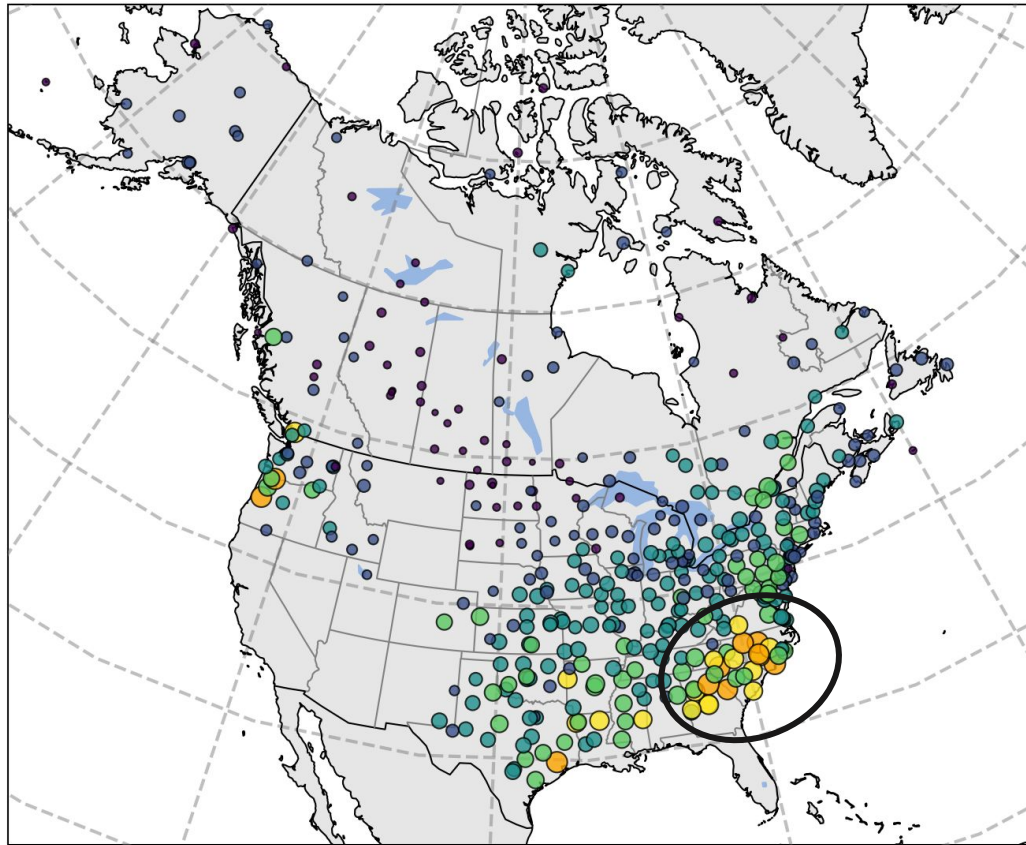


Median Annual Long-Duration (LD) Events

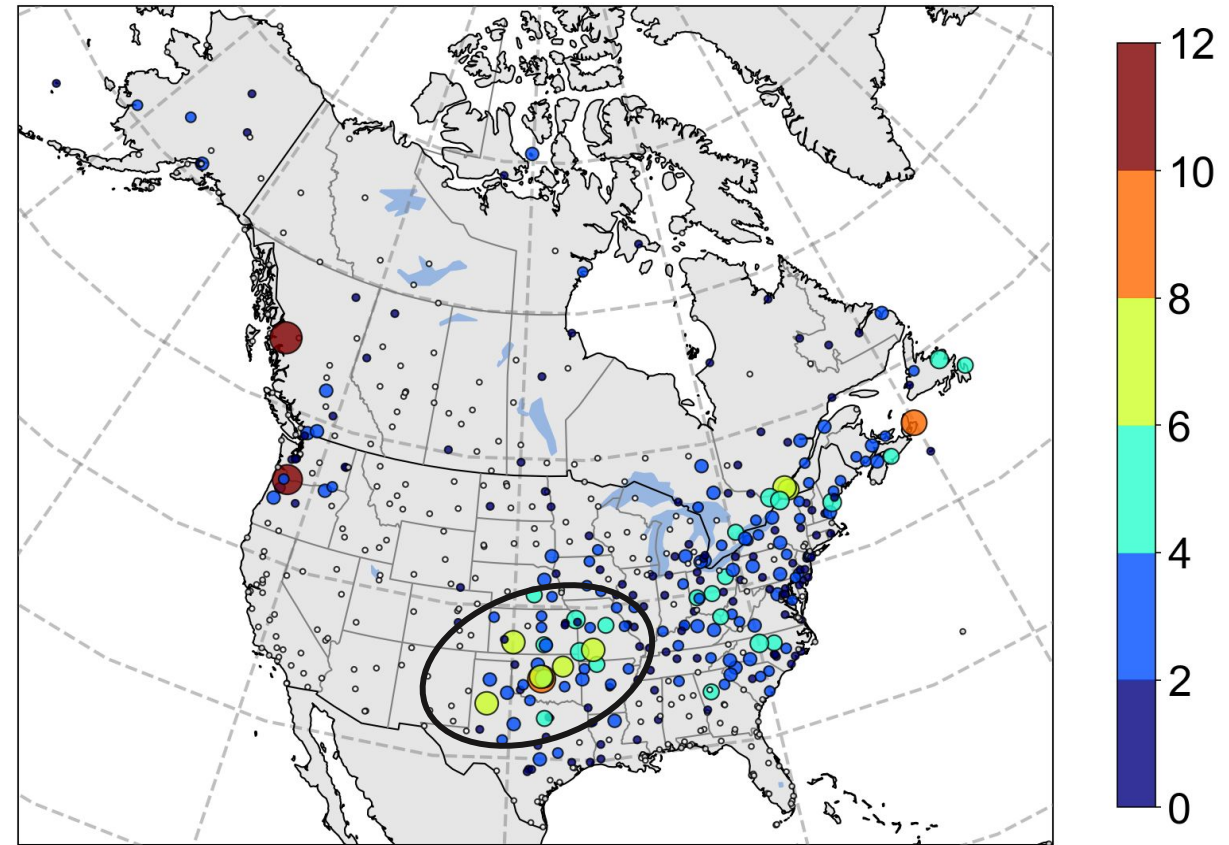


...but when freezing rain occurs in the Southeast and South Central U.S., it tends to be persistent

% of FZRA events that are long-duration

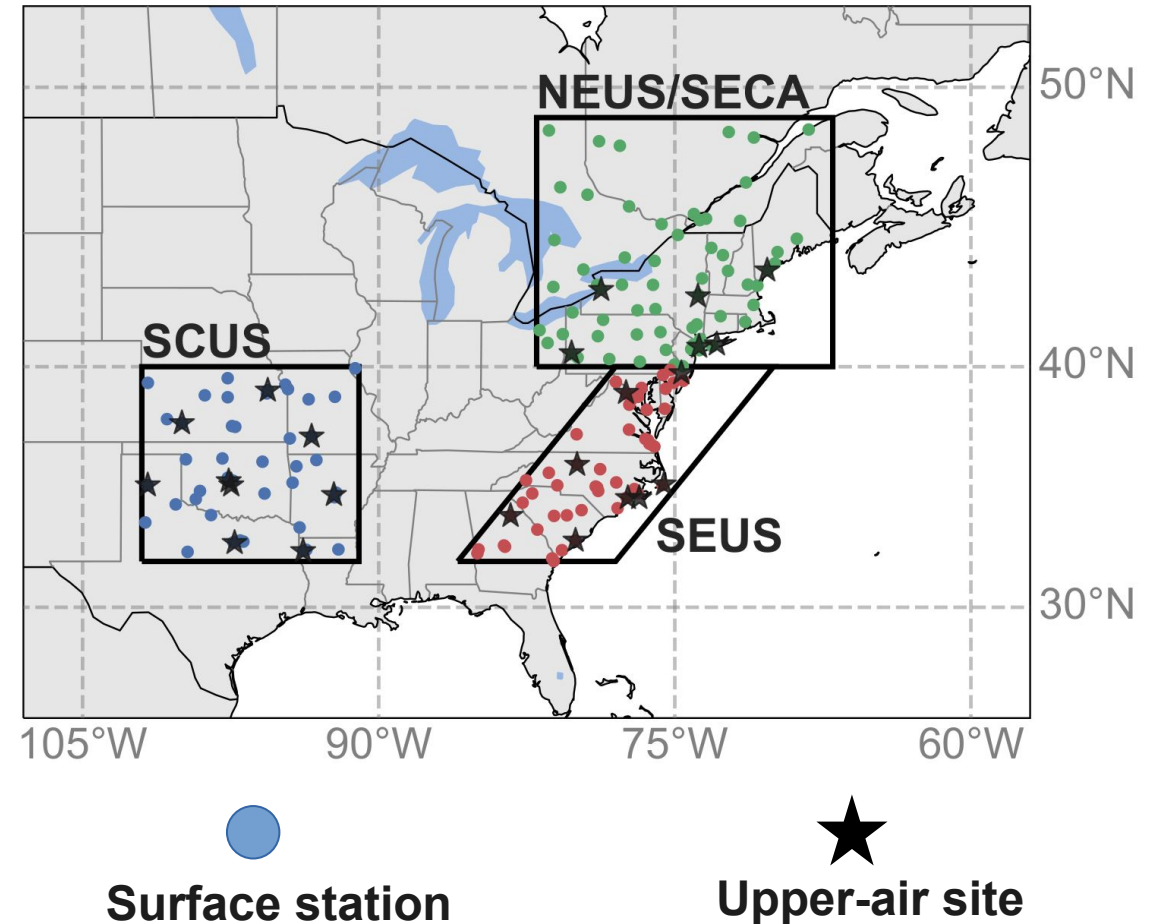


Number of 18+ h FZRA events (1979-2016)



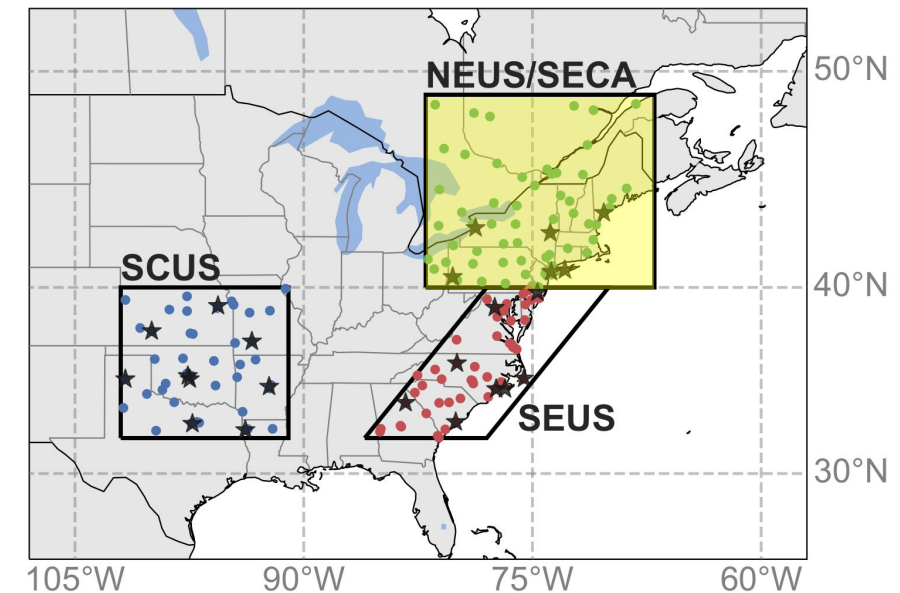
We identify three focus regions based on this climatology and examine FZRA events in each

- Northeastern U.S./ Southeastern Canada (**NEUS/SECA**)
- South Central U.S. (**SCUS**)
- Southeastern U.S. (**SEUS**)
- Upper-air observations
 - Regional aggregation
 - Examine all soundings for events that started **within 1 h** of a radiosonde release



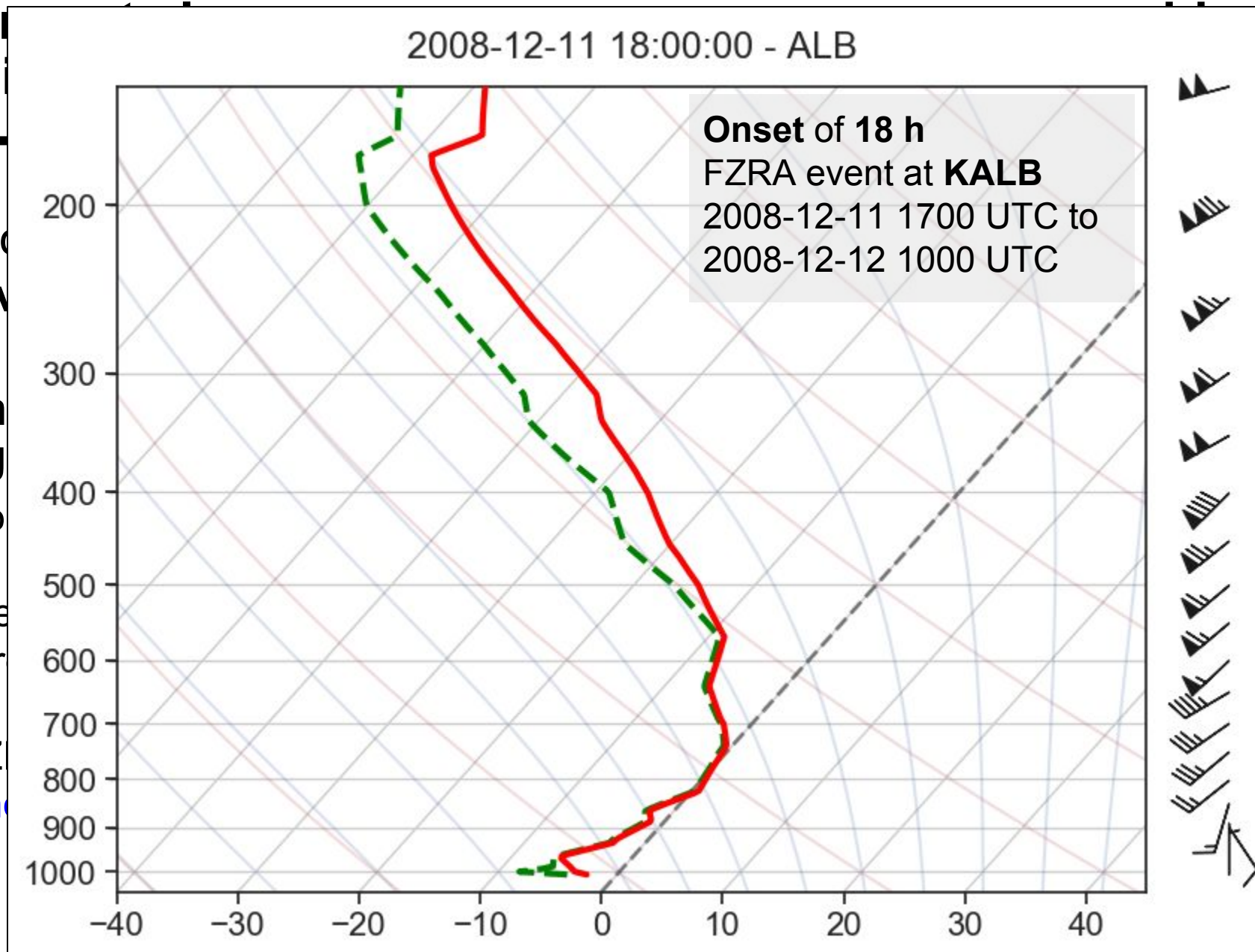
In McCray et al. (submitted to *Weather and Forecasting*), we identified the regional thermodynamic evolution of LD events

- We compared surface/upper-air obs taken at **LD event onset** with those at event **end**
- Northeastern U.S./ Southeastern Canada (**NEUS/SECA**)
 - Cold onset surface temperature, deep cold layer
 - Weak or absent surface cold-air advection
 - Weak onset warm layer
 - Strong warm-air advection just above the surface
 - Builds warm layer, erodes cold layer
 - FZRA ends as surface temperature reaches 0°C
 - Snow/Ice Pellets → FZRA → Rain

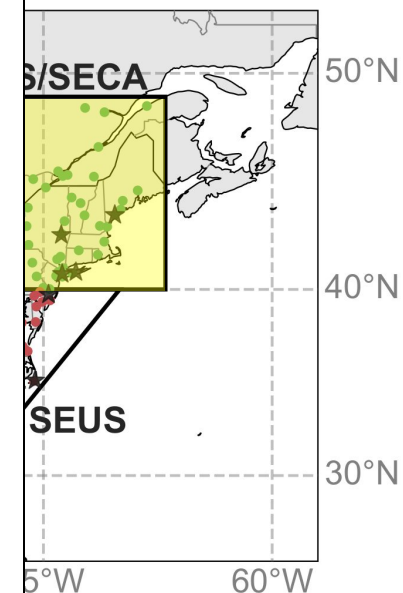


In McC
the reg

- We co
LD ev
- North
(NEU
 - Co
 -
 - We
 - Str
 -
 - FZ
 - Sn

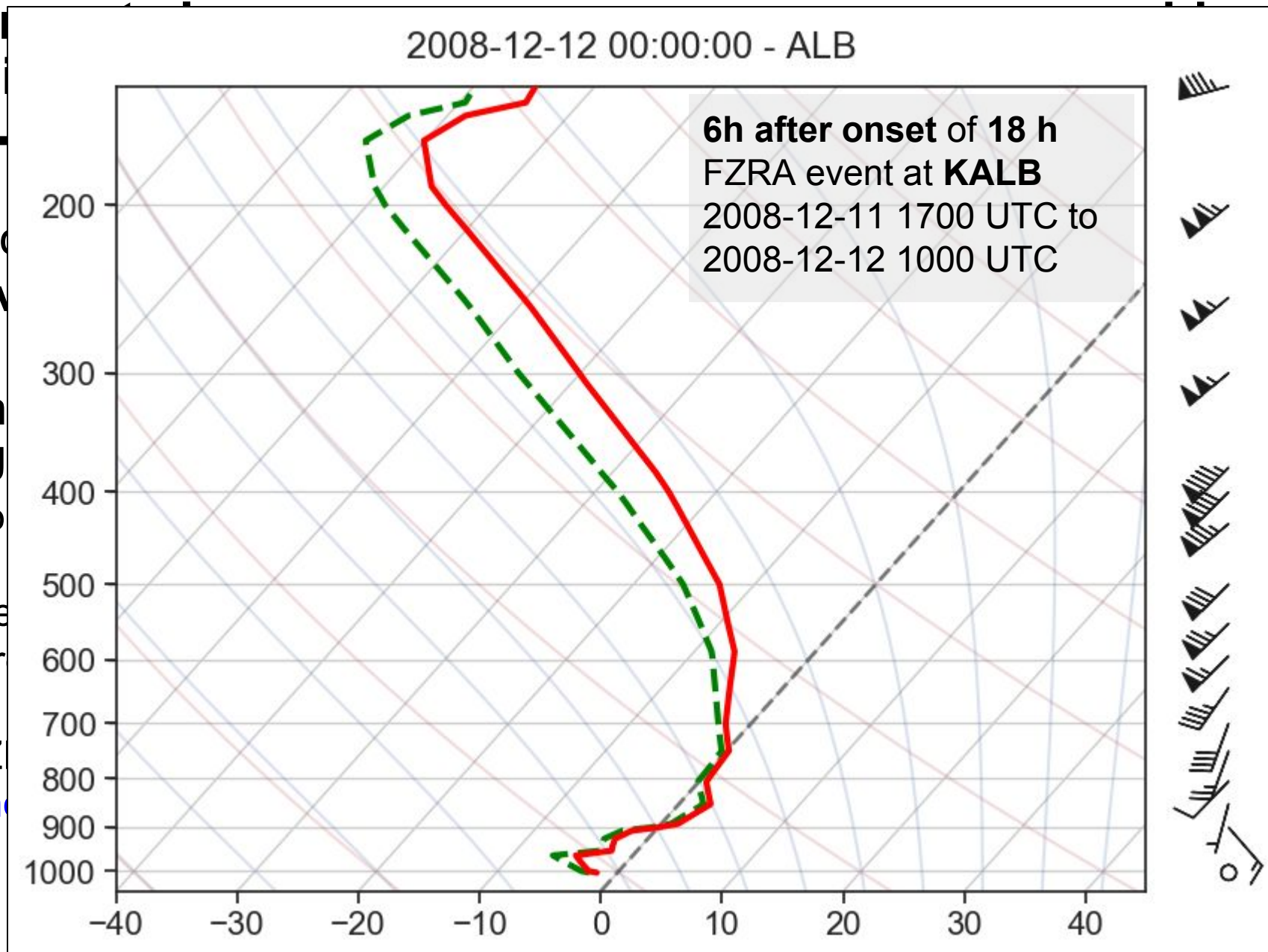


tified



In McC
the regi

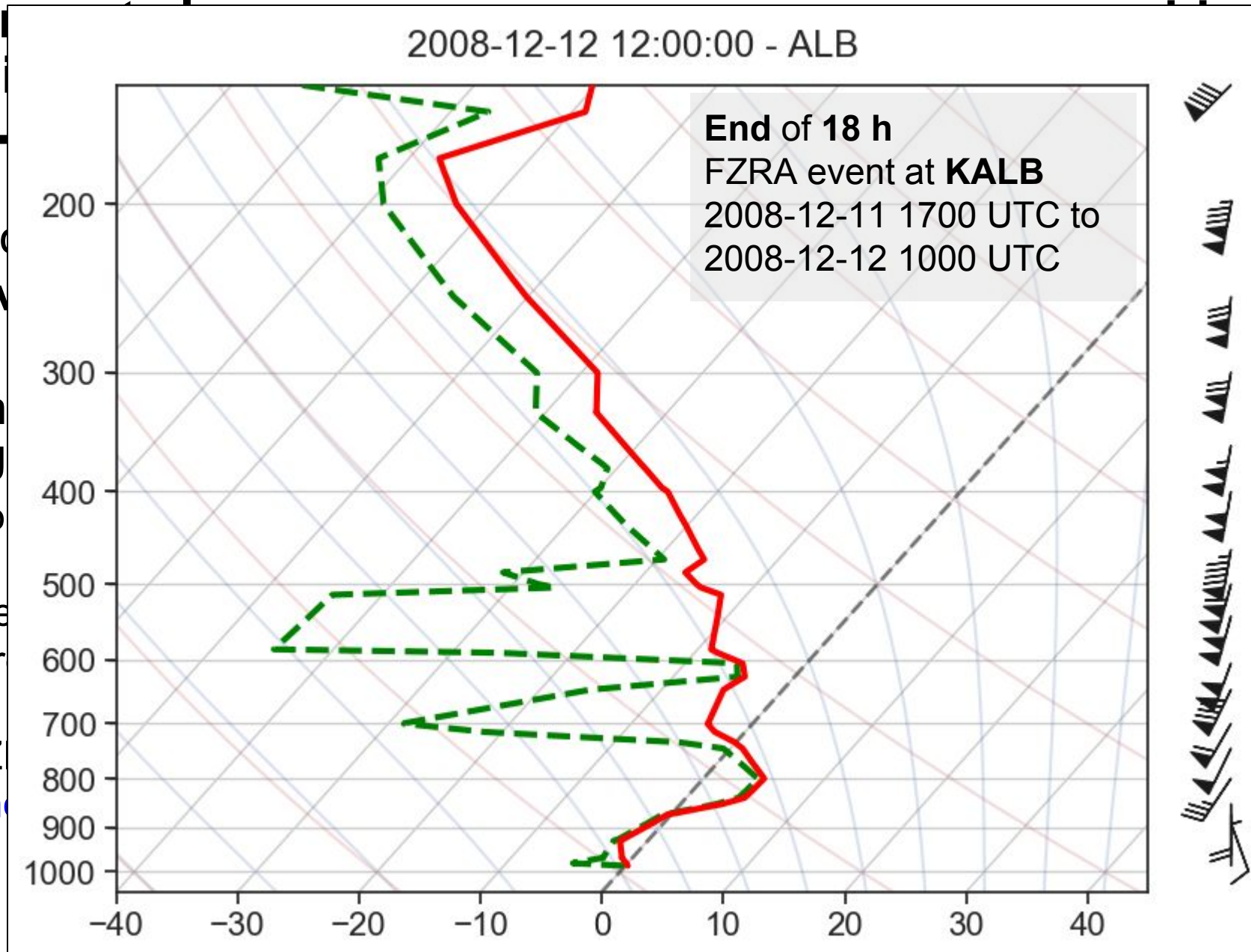
- We co
LD ev
- North
(NEU
 - Co
 -
 - We
 - Str
 -
 - FZ
 - Sn



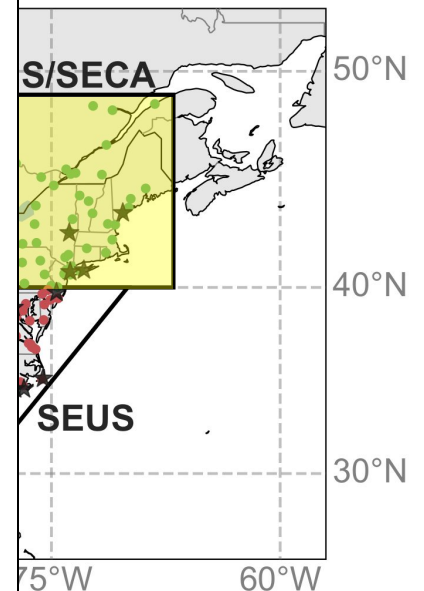
tified

In McC
the regi

- We co
LD ev
- North
(NEU
 - Co
 -
 - We
 - Str
 -
 - FZ
 - Sn



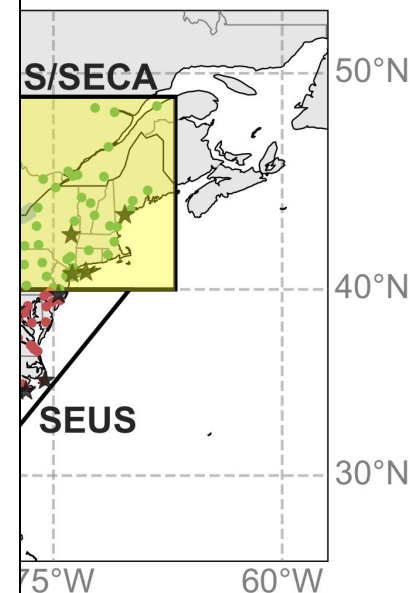
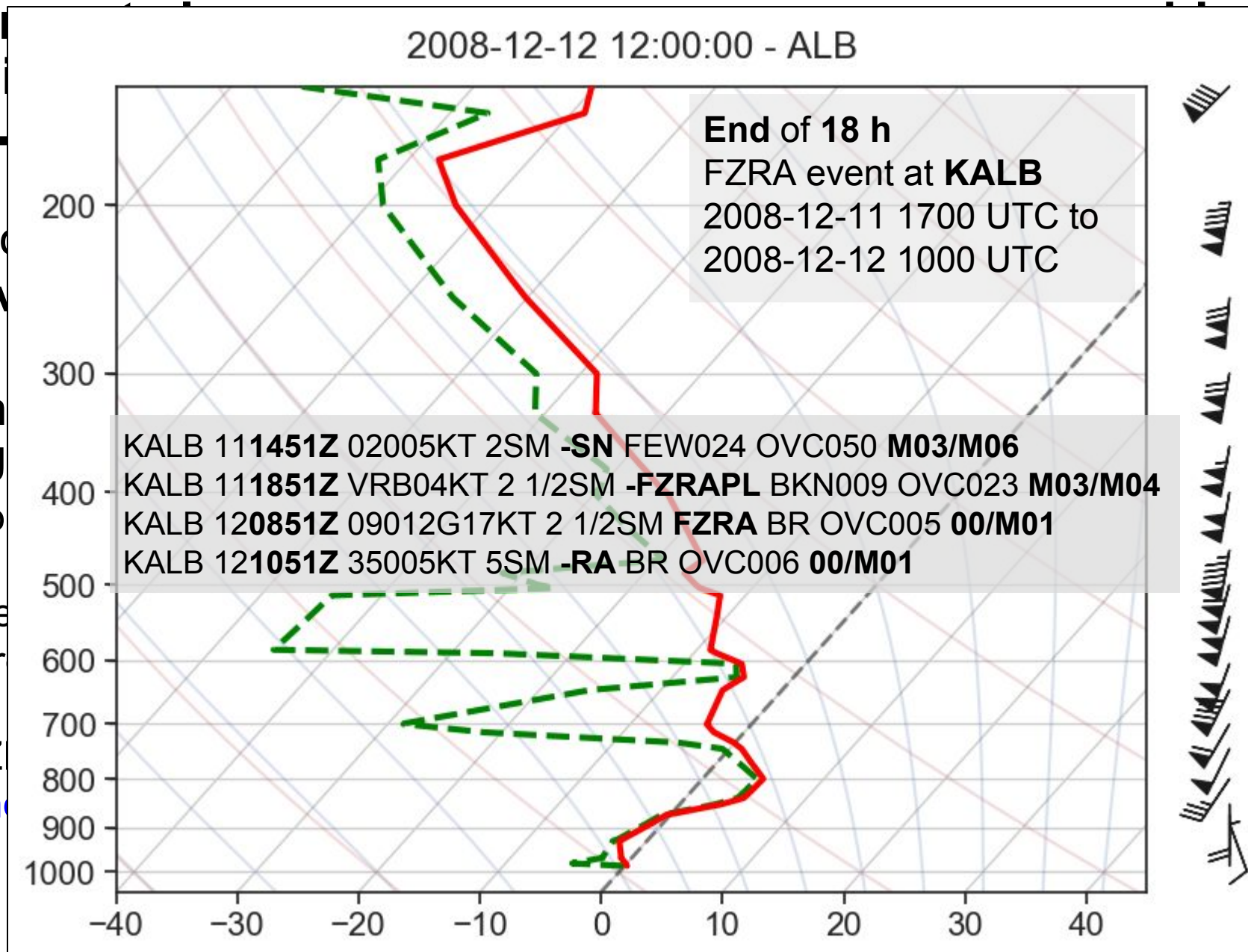
ified



In McC
the regi

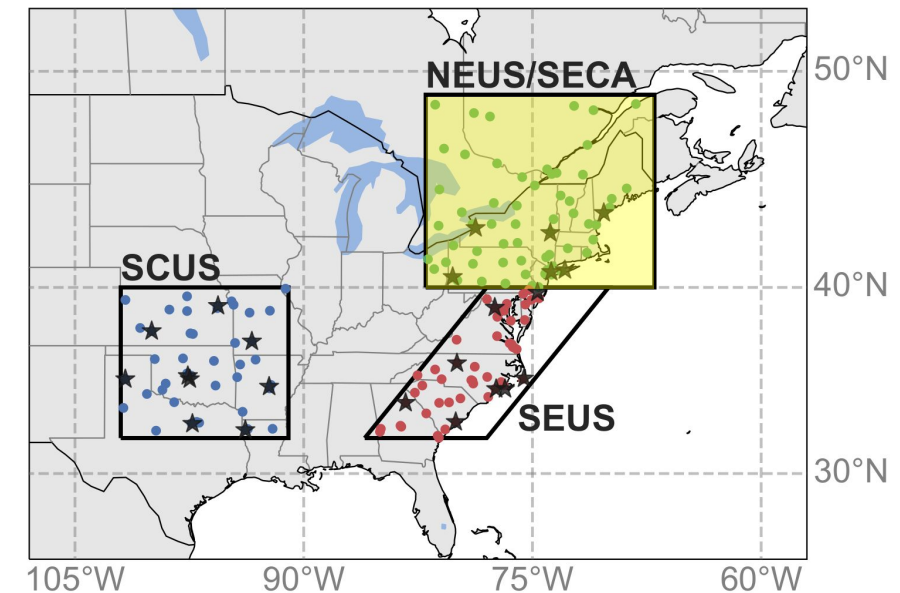
tified

- We co
LD ev
- North
(NEU
 - Co
 -
 - We
 - Str
 -
 - FZ
 - Sn



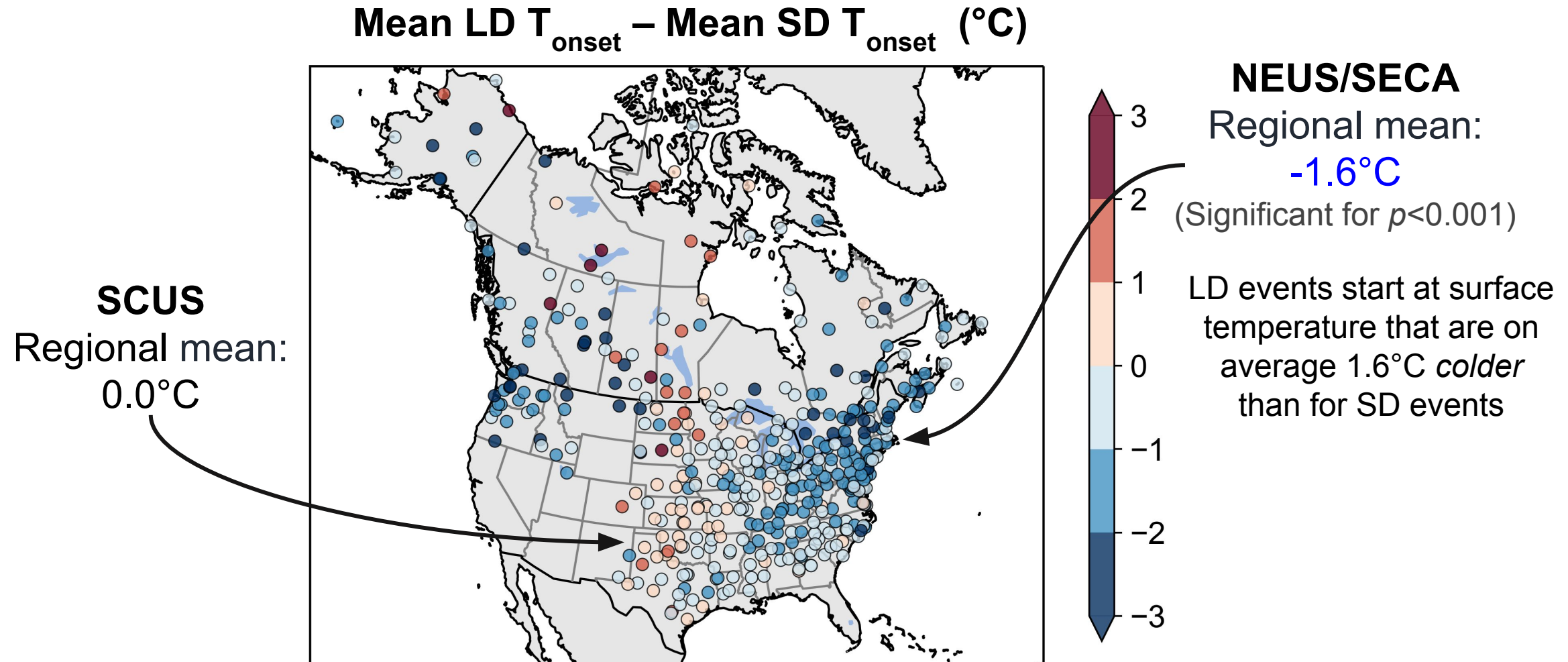
In McCray et al. (submitted to *Weather and Forecasting*), we identified the regional thermodynamic evolution of LD events

- We compared surface/upper-air obs taken at **LD event onset** with those at event **end**
- Northeastern U.S./ Southeastern Canada (**NEUS/SECA**)
 - Cold onset surface temperature, deep cold layer
 - Weak or absent surface cold-air advection
 - Weak onset warm layer
 - Strong warm-air advection just above the surface
 - Builds warm layer, erodes cold layer
 - FZRA ends as surface temperature reaches 0°C
 - Snow/Ice Pellets → FZRA → Rain

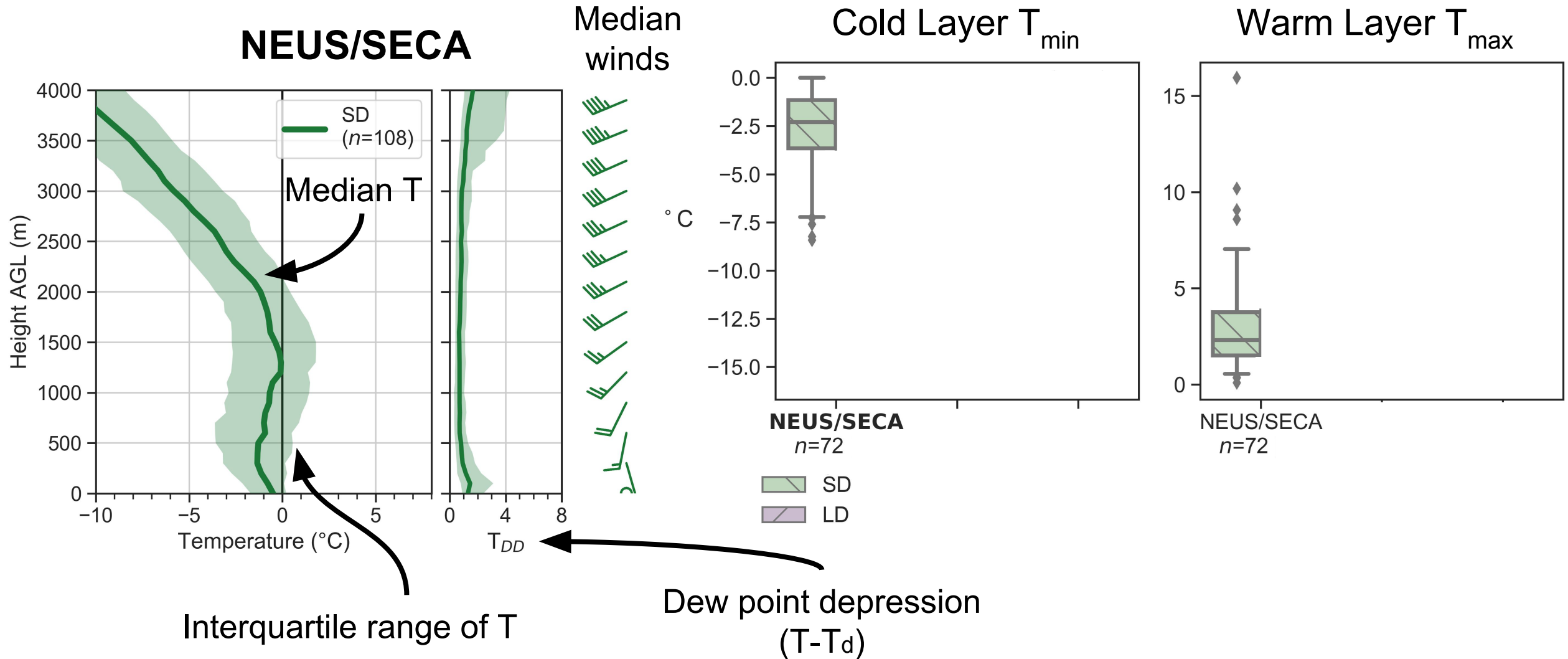


We now compare conditions at LD (long-duration) event onset with those at SD (short-duration) event onset

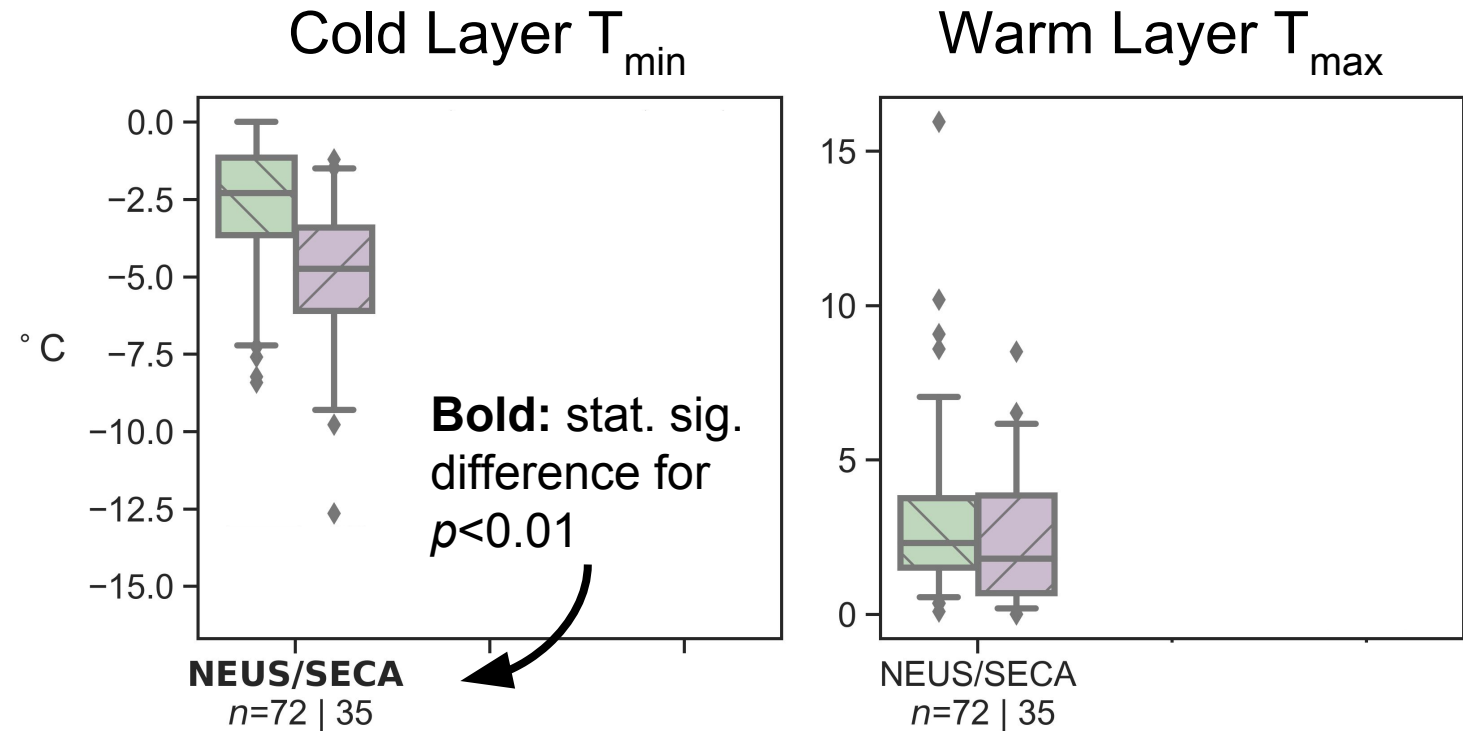
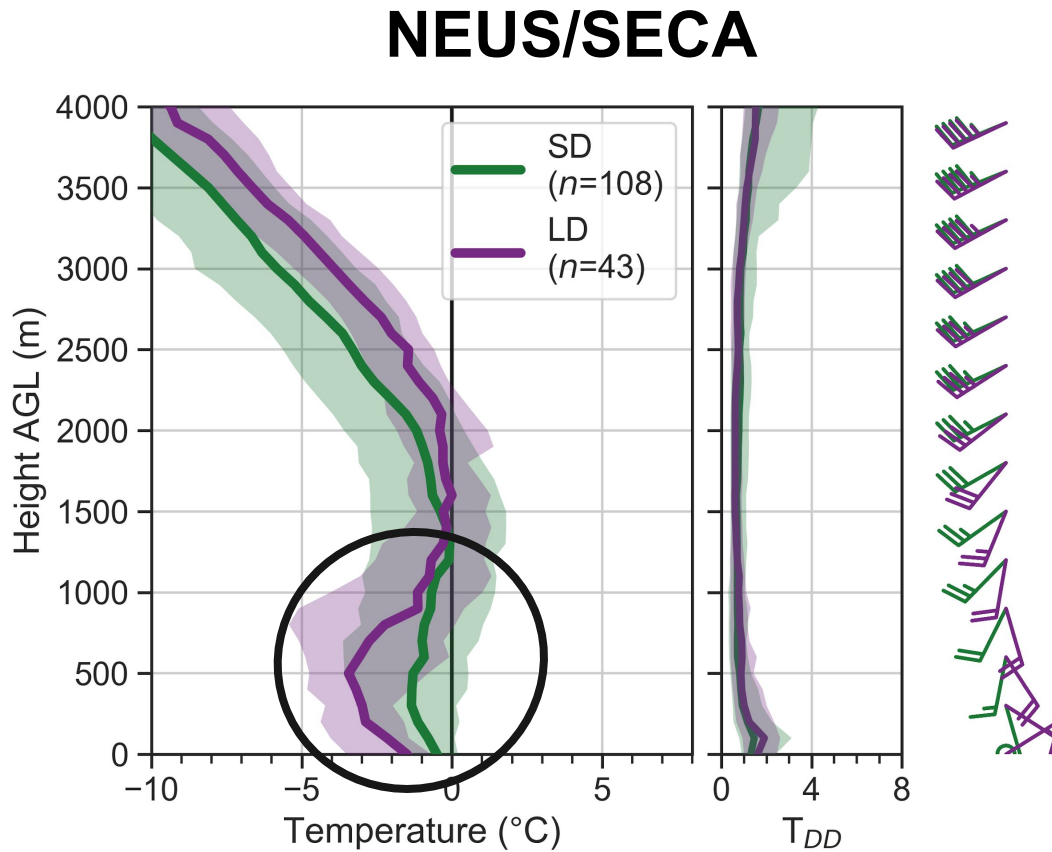
How do surface temperatures differ between long-duration (LD) and short-duration (SD) events?



How do thermodynamic profiles differ between SD and LD events?



A deeper and colder cold layer supports longer duration events over the NEUS/SECA

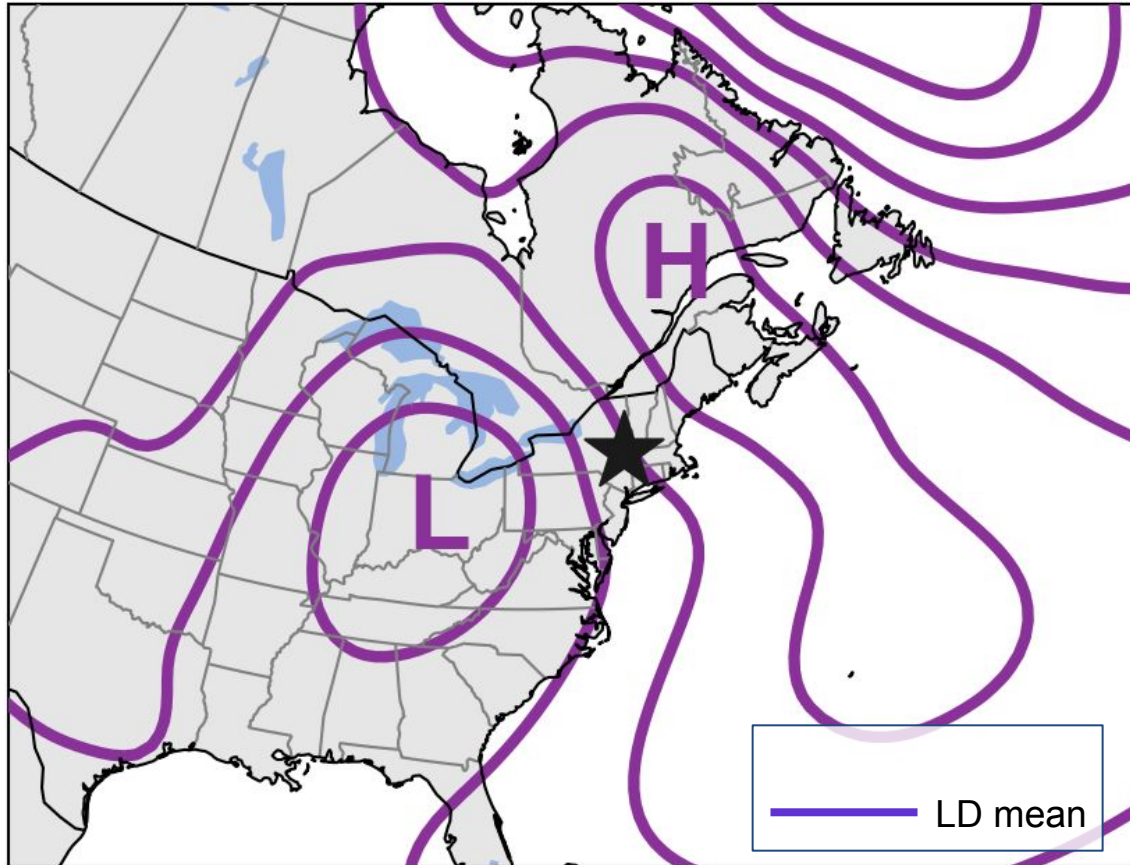


Median **cold layer** is **deeper** and **colder** at onset of LD events than SD events

→ Allows warming from freezing + advection to persist for longer period before layer surface warms to 0°C

How do synoptic patterns differ between SD and LD events?

Composite SLP at onset

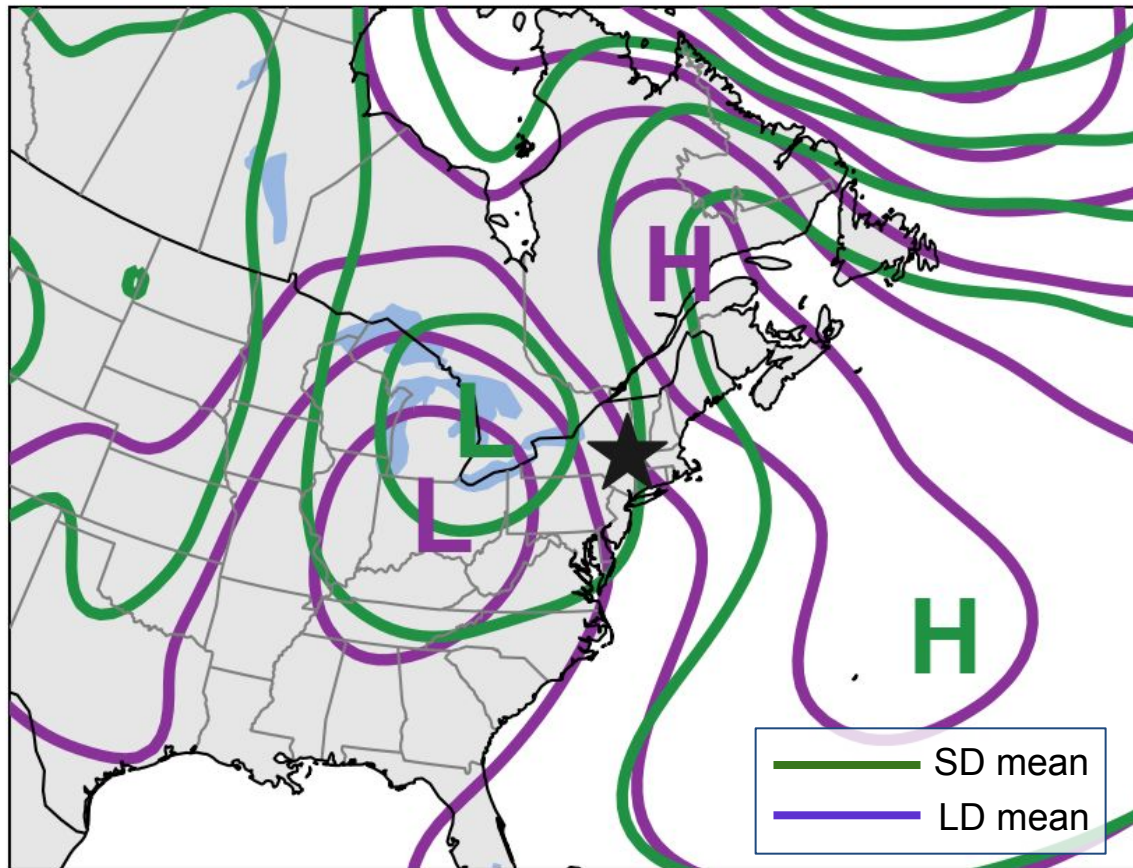


KALB - Albany, NY

- We composite a random sample of **30 LD** and **30 SD** events at select locations
- At KALB, LD events involve:
 - A cyclone situated to the SW
 - An anticyclone centered to the NE

How do synoptic patterns differ between SD and LD events?

Composite SLP at onset

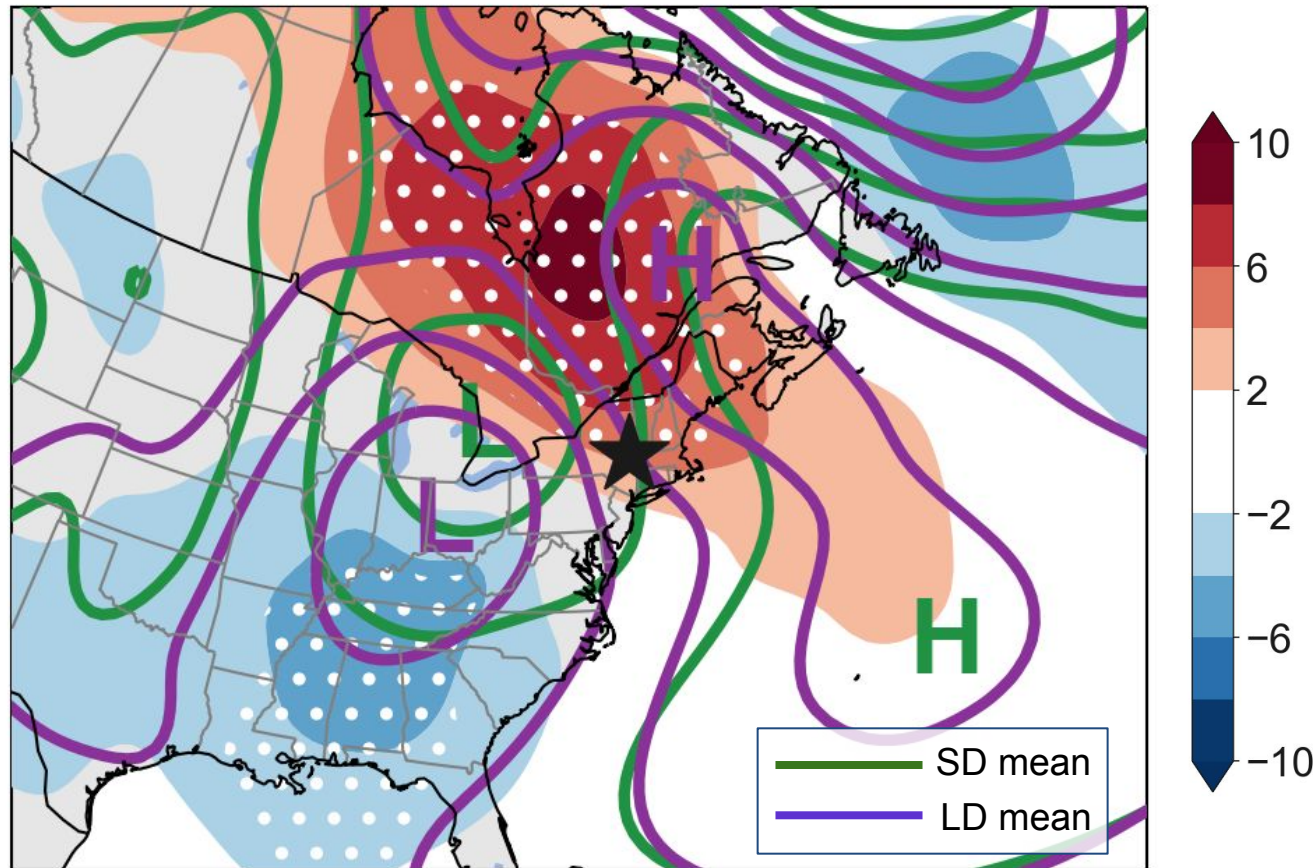


KALB - Albany, NY

- We composite a random sample of **30 LD** and **30 SD** events at select locations
- At KALB, LD events involve:
 - A cyclone situated to the SW
 - An anticyclone centered to the NE
- SD events involve
 - A cyclone situated due W
 - An anticyclone centered due E

How do synoptic patterns differ between SD and LD events?

LD mean SLP – SD mean SLP



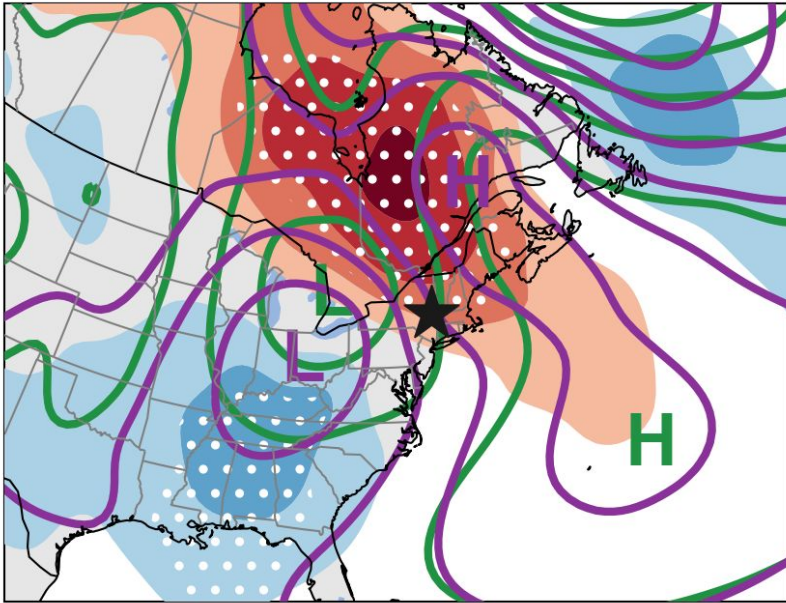
Stippling:
Significant for $p < 0.05$

KALB - Albany, NY

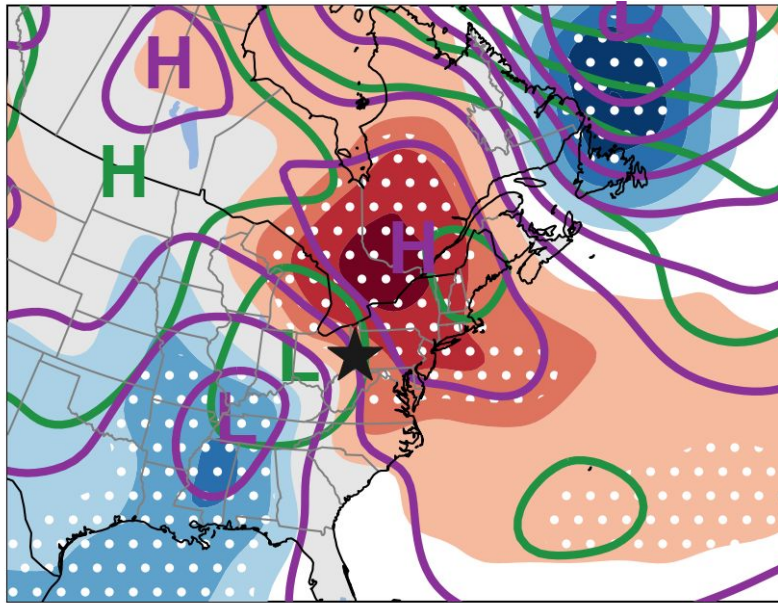
- We composite a random sample of **30 LD** and **30 SD** events at select locations
- At KALB, LD events involve:
 - A cyclone situated to the SW
 - An anticyclone centered to the NE
- SD events involve
 - A cyclone situated due W
 - An anticyclone centered due E

These differences are consistent within focus regions, despite terrain variations, etc.

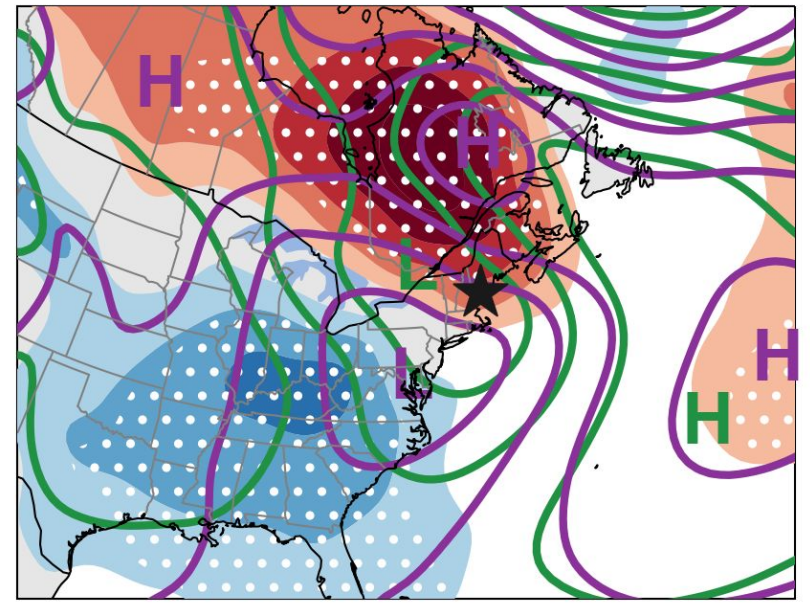
KALB - Albany, NY



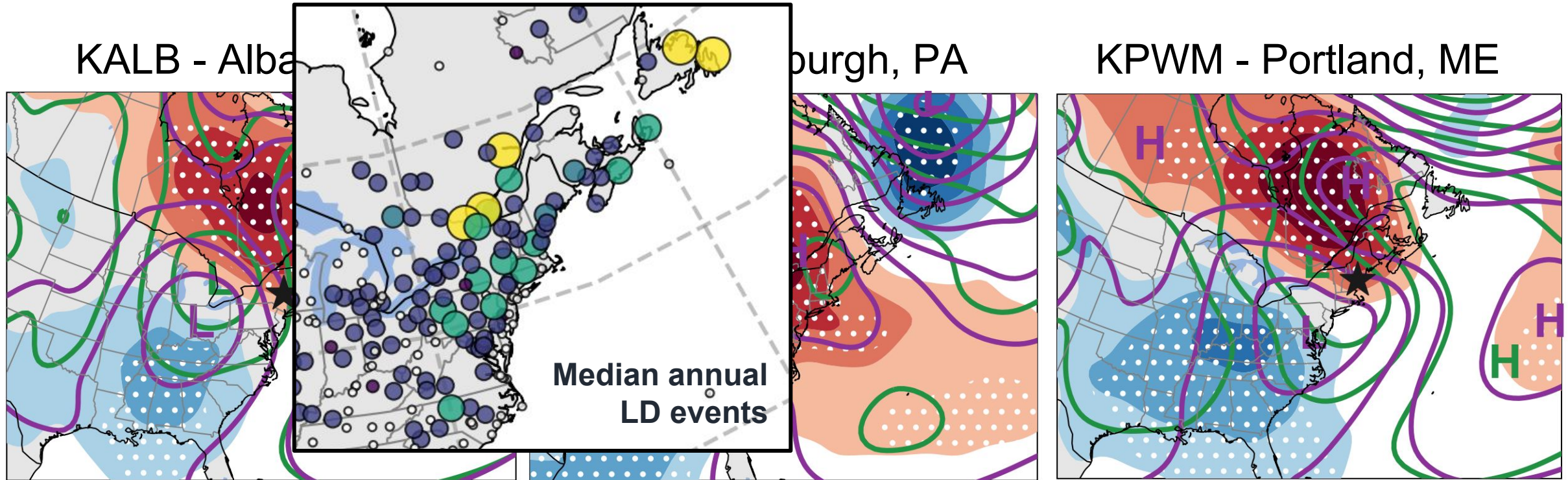
KPIT - Pittsburgh, PA



KPWM - Portland, ME



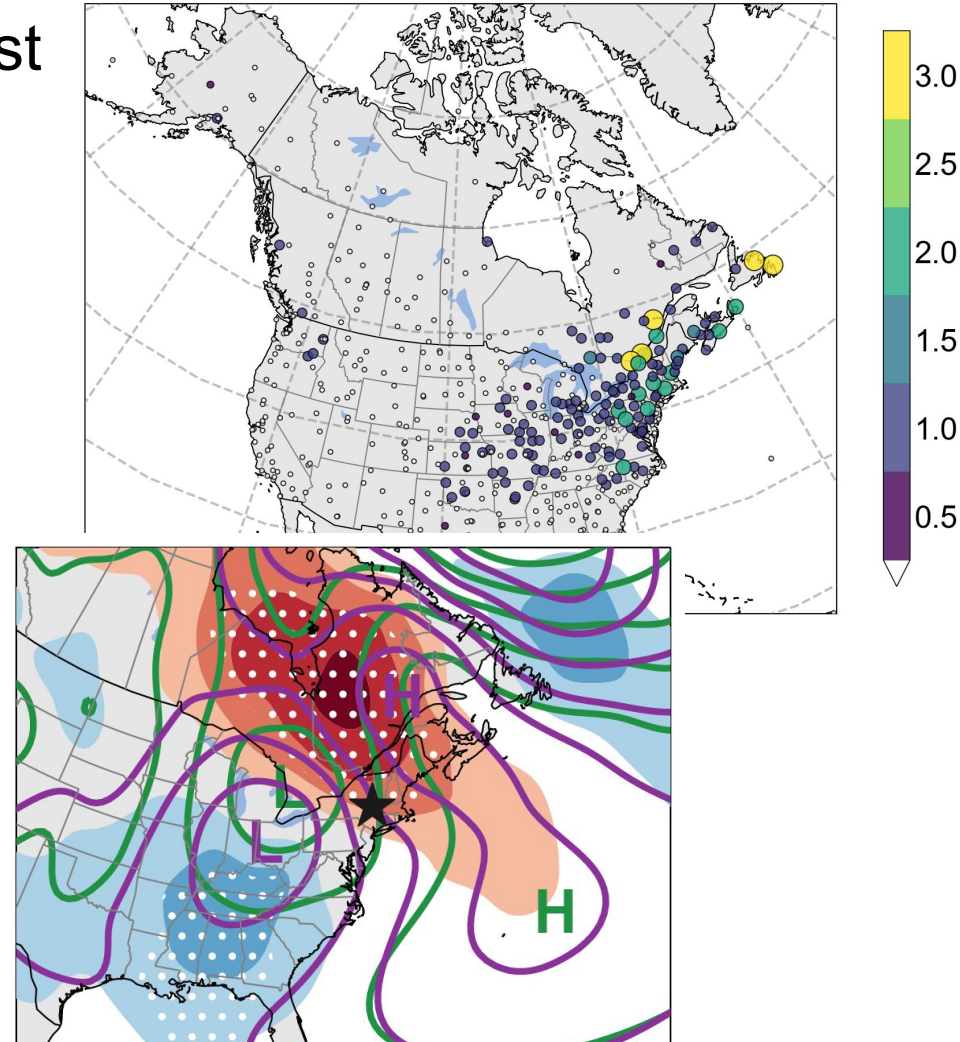
These differences are consistent within focus regions, despite terrain variations, etc.



Terrain features can support **cold air trapping/channelling** during favorable synoptic setups for freezing rain
→ Allows for longer duration events than in areas without such features

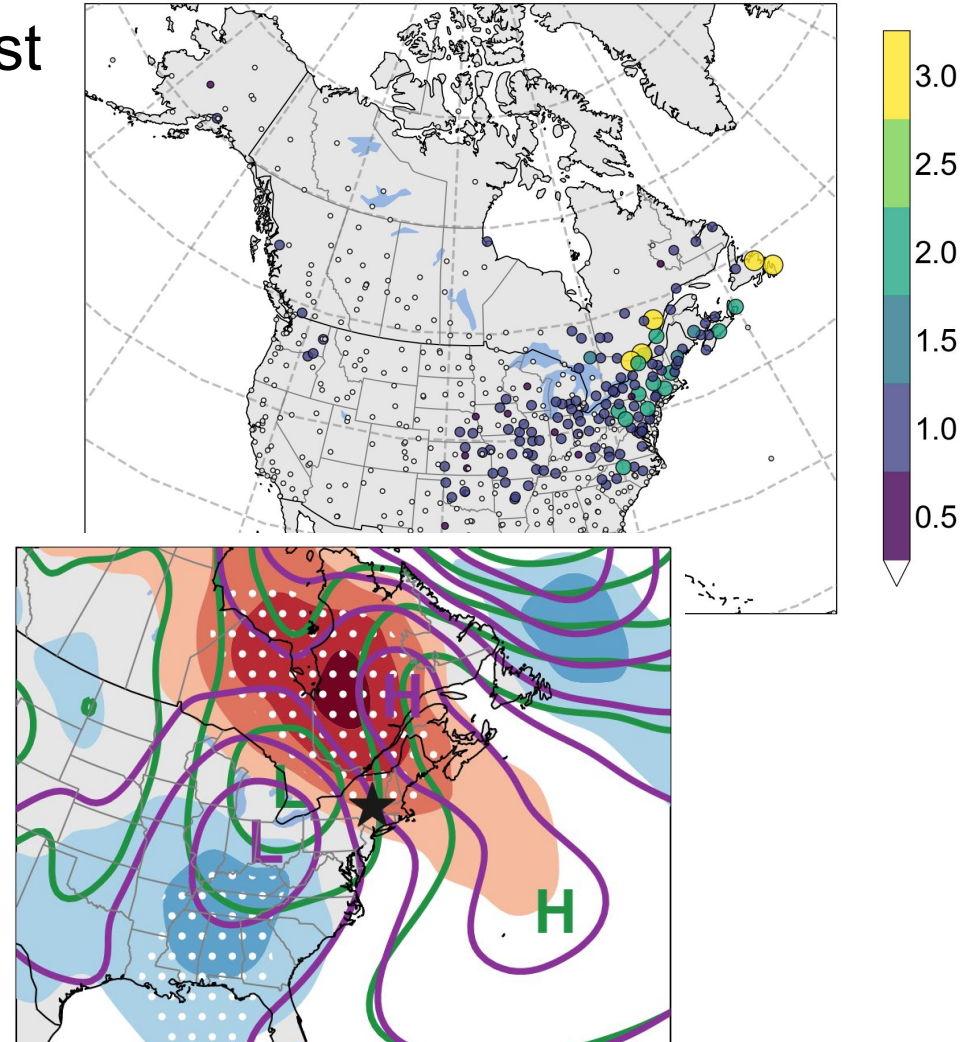
Summary

- Long-duration (LD) freezing rain events are most common over the **northeastern U.S. and southeastern Canada**
 - Storm track key, terrain produces local maxima
- **LD events**, compared with SD ones, exhibit
 - A deeper, colder onset cold layer
 - Colder surface temperatures at onset
 - A deeper anticyclone to the NE (instead of E)
 - A more distant surface cyclone/warm front



Summary

- Long-duration (LD) freezing rain events are most common over the **northeastern U.S. and southeastern Canada**
 - Storm track key, terrain produces local maxima
- **LD events**, compared with SD ones, exhibit
 - A deeper, colder onset cold layer
 - Colder surface temperatures at onset
 - A deeper anticyclone to the NE (instead of E)
 - A more distant surface cyclone/warm front
- **Future Work**
 - Further analysis of composite differences
 - How can relationships between onset characteristics and duration be applied to forecasts?



References

- Cortinas, J. V., B. C. Bernstein, C. C. Robbins, J. Walter Strapp, J. V. Cortinas Jr., B. C. Bernstein, C. C. Robbins, and J. Walter Strapp, 2004: An Analysis of Freezing Rain, Freezing Drizzle, and Ice Pellets across the United States and Canada: 1976–90. *Wea. Forecasting*, **19**, 377–39.
- Kain, J. S., S. M. Goss, and M. E. Baldwin, 2000: The Melting Effect as a Factor in Precipitation-Type Forecasting. *Wea. Forecasting*, **15**, 700–714.
- Lackmann, G. M., K. Keeter, L. G. Lee, and M. B. Ek, 2002: Model Representation of Freezing and Melting Precipitation: Implications for Winter Weather Forecasting. *Wea. Forecasting*, **17**, 1016–1033.
- McCray, C. D., E. H. Atallah, and J. R. Gyakum: Long-Duration Freezing Rain Events over North America: Regional Climatology and Thermodynamic Evolution. Submitted to *Wea. Forecasting*.
- Stewart, R. E., 1985: Precipitation types in winter storms. *Pure Appl. Geophys.*, **123**, 597–609.